Lecture: 13		GY
Topics Cove	ered: CLIMATOLOGY	
Atmosphere	1	
• Comp	osition: Gases, Water Vapor, Dust Part	icles
Struct	ture: Exosphere, Thermosphere, Mesos	phere, Stratosphere, Troposphere,
• Temp	erature: Altitude vs. Temperature, Inve actors controlling Temperature distribu	ersion of Temperature
 Distri 	bution of Temperature Month of Janua	iry-July
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Climatology		
Climatology	is the colontific study of climate	It is a branch of the atmosphe
sciences concer	ned with the description of the clin	nate, the analysis of the causes
climatic differe r	ces and changes and their practical	consequences.
Climate vs W	eather	
Parameter	Climate	Weather
Definition	Describes the average conditions expected at a specific place at a given time. A region's climate is	Describes the atmospheric conditions at a specific place at a specific point in time. Weather
	generated by the climate system, which has five components: atmosphere, hydrosphere, cryosphere, land surface, and	generally refers to day-to-day temperature and precipitation activity
Components	Climate may include precipitation, temperature, humidity, sunshine,	Weather includes sunshine
	wind velocity, phenomena such as fog, frost, and hail storms over a long period of time.	rain, cloud cover, winds, hair snow, sleet, freezing rain flooding, blizzards, ice storms thunderstorms, steady rains from a cold front or warm front excessive heat, heat waves and
Forecast	wind velocity, phenomena such as fog, frost, and hail storms over a long period of time. By aggregates of weather statistics over periods of 30 years	 rain, cloud cover, winds, nail snow, sleet, freezing rain flooding, blizzards, ice storms thunderstorms, steady rains from a cold front or warm front excessive heat, heat waves and more By collecting meteorologica data, like air temperature pressure, humidity, solar radiation, wind speeds and direction etc.

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Parameter	Climate	Weather
About	Climate is defined as statistical weather information that describes the variation of weather at a given place for a specified interval.	Weather is the day-to-day state of the atmosphere, and its short-term (minutes to weeks) variation
Time period	Measured over a long period	Measured for short term
Study	Climatology	Meteorology

ATMOSPHERE

The atmosphere is the layer of gases that surrounds a celestial body, such as a planet or a moon.

- Mars, a neighboring planet to Earth, also has an atmosphere, although it is much thinner compared to Earth's. The Martian atmosphere is predominantly composed of carbon dioxide (around 95.3%), with nitrogen, argon, and trace amounts of other gases. The thin atmosphere on Mars cannot support human life as we know it, and it does not provide significant protection against harmful solar radiation.
- Venus, another neighboring planet, has an atmosphere that is mainly composed of carbon dioxide (around 96.5%) and thick clouds of sulfuric acid. The atmosphere on Venus is incredibly dense and generates a strong greenhouse effect, resulting in surface temperatures that can melt lead. This makes Venus the hottest planet in our solar system, despite being farther from the Sun than Mercury.
- On **Earth**, the atmosphere is crucial for supporting life and plays a vital role in regulating the planet's **climate**. The Earth's atmosphere is a dynamic and complex mixture of gases, water vapor, and dust particles. These components also play crucial roles in **shaping weather patterns**.
- The earth is enveloped entirely by air, creating a spherical shape that is flatter at the poles and bulging at the Equator due to the combined influence of gravity and the Earth's rotation, similar to an egg's shape.
- The atmosphere can be divided into several **layers** or **spheres**, each concentric with the **Earth**, exhibiting different **temperature** and **chemical characteristics**.
- The force of **gravity** is responsible for holding the atmosphere to the **planet**, and it also determines the composition of **gases** present within it.

Gases in the Atmosphere

- The atmosphere consists mainly of different gases, with nitrogen and oxygen being the most abundant. Other gases like carbon dioxide, water vapor, methane, ozone, and noble gases like argon are also present in smaller quantities.
- Nitrogen (N₂): It is the most abundant gas, making up approximately 78% of the air we breathe. Plants and animals use nitrogen to build proteins and other organic molecules.



≤Notes			
	 Oxygen (O₂): Oxygen according for respiration, as organis respiration. 	ounts for about 21% of the atmostrian of the atmostrian second se	sphere. It is esse food through ce
	 Carbon Dioxide (CO₂): Tho dioxide plays a crucial role warm enough to support have raised CO₂ levels, cor 	ugh it constitutes a small fraction (a in the greenhouse effect. It traps he life. Increased human activities lik atributing to global warming.	round 0.04%), ca eat, keeping the f e burning fossil
	 Methane (CH₄): Methane i is relatively low. It is rele livestock, and fossil fuel ex 	s a potent greenhouse gas, even tho ased from various sources, includi traction.	ough its concentra ing natural wetla
	Ozone (O ₃): Ozone is a m	olecule with three oxygen atoms.	It exists in two
	layers of the atmosphere harmful ultraviolet (UV) r pollution and smog format	: the stratosphere, where it shield adiation, and the troposphere, wh tion. Gas	ds life on Earth here it contributo Percentage
	layers of the atmosphere harmful ultraviolet (UV) r pollution and smog formation	: the stratosphere, where it shield adiation, and the troposphere, wh tion. Gas	ds life on Earth nere it contributo Percentage by Volume
	layers of the atmosphere harmful ultraviolet (UV) r pollution and smog forma All other gases	: the stratosphere, where it shield adiation, and the troposphere, wh tion. Gas Nitrogen (N ₂)	Percentage by Volume 78.084
	layers of the atmosphere harmful ultraviolet (UV) r pollution and smog format All other gases	: the stratosphere, where it shield adiation, and the troposphere, wh tion. Gas Nitrogen (N ₂) Oxygen (O ₂)	Percentage by Volume 78.084 20.946
	layers of the atmosphere harmful ultraviolet (UV) r pollution and smog format All other gases 1% Oxygen	: the stratosphere, where it shield adiation, and the troposphere, wh tion. Gas Nitrogen (N ₂) Oxygen (O ₂) Argon (Ar)	Percentage by Volume 78.084 20.946 0.934
	layers of the atmosphere harmful ultraviolet (UV) r pollution and smog format All other gases 1% Oxygen 21%	: the stratosphere, where it shield adiation, and the troposphere, wh tion. Gas Nitrogen (N ₂) Oxygen (O ₂) Argon (Ar) Carbon dioxide (CO ₂)	Percentage by Volume 78.084 20.946 0.934 0.037
	layers of the atmosphere harmful ultraviolet (UV) r pollution and smog format All other gases	: the stratosphere, where it shield adiation, and the troposphere, wh tion. Gas Nitrogen (N ₂) Oxygen (O ₂) Argon (Ar) Carbon dioxide (CO ₂) Neon (Ne)	Percentage by Volume 78.084 20.946 0.934 0.037 0.00182
	layers of the atmosphere harmful ultraviolet (UV) r pollution and smog formar All other gases 1% Oxygen 21% Nitrogen 78%	: the stratosphere, where it shield adiation, and the troposphere, wh tion. Gas Nitrogen (N ₂) Oxygen (O ₂) Argon (Ar) Carbon dioxide (CO ₂) Neon (Ne) Helium (He)	Percentage by Volume 78.084 20.946 0.934 0.037 0.00182 0.00052
	layers of the atmosphere harmful ultraviolet (UV) r pollution and smog formar All other gases 1% Oxygen 21% Nitrogen 78%	: the stratosphere, where it shield adiation, and the troposphere, wh tion. Gas Nitrogen (N ₂) Oxygen (O ₂) Argon (Ar) Carbon dioxide (CO ₂) Neon (Ne) Helium (He) Methane (CH ₄)	Percentage by Volume 78.084 20.946 0.934 0.037 0.00182 0.00052 0.00015

Water Vapor

Water vapor is the gaseous form of water in the atmosphere. It plays a crucial role in the **water cycle**, evaporating from water bodies, forming clouds, and eventually falling as rain or snow. Water vapor influences weather patterns, cloud formation, and precipitation.

Examples of Weather Phenomena Involving Water Vapor:

- Evaporation and Condensation: Water bodies exposed to sunlight undergo evaporation, converting water into water vapor in the air. When the air cools, water vapor condenses into clouds.
- **Humidity:** Humidity represents the amount of water vapor in the air relative to its maximum capacity. High humidity can make us feel uncomfortable.
- **Clouds:** Rising moist air cools and forms clouds as water vapor condenses into tiny droplets or ice crystals.
- **Precipitation:** Water droplets or ice crystals in clouds grow and fall to the Earth as rain, snow, sleet, or hail.
- **Greenhouse Effect:** Water vapor is a **potent greenhouse gas** that traps heat in the Earth's atmosphere, maintaining a habitable temperature range.

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Dust Particles

Dust particles in the atmosphere are tiny solid particles originating from natural sources like wind erosion, volcanic eruptions, and dust storms, as well as human activities like burning fossil fuels, agriculture, and construction.

Examples of Dust Particles:

- Mineral Dust: Commonly from deserts and dry regions, carried by wind.
- Volcanic Ash: Released during volcanic eruptions, affecting weather and posing hazards.
- **Soot and Black Carbon:** Resulting from incomplete combustion, contributing to air pollution and global warming.
- **Pollen:** Airborne during pollen season, causing allergies in some individuals.

Effects of Dust Particles in the Atmosphere:

- Atmospheric Cooling: Certain dust particles, like volcanic ash, reflect sunlight, causing temporary cooling.
- Air Quality and Health: Inhalation of dust particles can lead to respiratory problems.
- **Climate Impact:** Dust particles influence cloud formation and weather patterns.
- **Nutrient Transport:** Dust can carry essential nutrients like phosphorus and iron to other ecosystems.

Understanding atmospheric composition is vital for addressing environmental challenges and ensuring a sustainable future.

Structure of Atmosphere

Structure of Atmosphere

The Earth's atmosphere is divided into several distinct layers based on their temperature and composition. Each layer has its **unique characteristics** and plays a **crucial role** in the Earth's climate and **atmospheric processes**. Let's explore the **structure of the atmosphere** and the features of each layer:

Troposphere:

- The troposphere is the lowest layer of the atmosphere, extending from the Earth's surface up to an average altitude of about 7 to 17 kilometers (4 to 11 miles).
- Features:
 - Weather: The troposphere is where most of the weather phenomena occur. As air rises and cools, it can condense to form clouds and precipitation. Vertical air movements in the troposphere lead to the formation of thunderstorms, hurricanes, and other weather events.
 - Jet Streams: High-altitude, fast-moving air currents known as jet streams are found within the upper part of the troposphere, influencing weather patterns and air travel routes.
- **Boundary:** The upper boundary of the troposphere is called the **tropopause**. It acts as a stable boundary between the troposphere and the stratosphere and varies in height with latitude and season.
- For Example:
 - **O** Rainfall occurs in the troposphere as warm, moist air rises and cools, leading to **condensation** and **cloud formation**.

<i>⊯Notes</i>	
	• Thunderstorms and hurricanes , driven by convective processes , are typical weather events that occur within the troposphere.
Strato	osphere:
•	The stratosphere lies above the troposphere and extends up to about 50 kilometers (31 miles) above the Earth's surface.
•	Features:
	\bigcirc Ozone Layer: The stratosphere contains a high concentration of ozone (O ₃) in the ozone layer. Ozone absorbs most of the Sun's harmful UV radiation, protecting life on Earth from its damaging effects.
•	Boundary: The upper boundary of the stratosphere is called the stratopause . It separates the stratosphere from the mesosphere and marks the region where the temperature starts to increase with altitude.
•	For Example: Stratospheric ozone is crucial for protecting life on Earth from harmful UV radiation. Human activities, such as the release of ozone-depleting substances , can lead to thinning of the ozone layer, resulting in increased UV radiation reaching the Earth's surface.
Meso	sphere:
•	The mesosphere is located above the stratosphere and extends up to about 85 kilometers (53 miles) in altitude.
•	Features:
	• Shooting Stars: The mesosphere is where most meteoroids burn up upon entering the Earth's atmosphere. The friction with the air causes them to heat up and produce streaks of light known as shooting stars or meteors.
•	Boundary: The upper boundary of the mesosphere is called the mesopause. It separates the mesosphere from the thermosphere and is characterized by its extremely low temperatures.
•	For Example: No specific human activities or technology are directly related to the mesosphere. However, it plays a role in protecting the Earth's surface from meteoroid impacts.
Therr	nosnhere
•	The thermosphere lies above the mesosphere and extends up to about 600 kilometers (372 miles) above the Earth's surface.
•	Features:
	• Auroras: The thermosphere is where the auroras, such as the Northern and Southern Lights, occur. Charged particles from the solar wind interact with gases in the thermosphere, producing the beautiful light displays.
•	Boundary: The upper boundary of the thermosphere is not precisely defined, as the transition between the thermosphere and the exosphere is gradual. This region is called the thermopause .
•	Example: The International Space Station (ISS) orbits within the lower thermosphere. Its relatively low atmospheric density allows it to stay in orbit while experiencing very little drag.
lonos	phere:
•	The ionosphere is a region within the thermosphere and the upper mesosphere. It is characterized by the presence of ionized gases (ions) due to the absorption of solar ultraviolet and X-ray radiation.



• The exosphere is the outermost layer of the Earth's atmosphere, transitioning into outer space. It extends beyond the thermosphere and has a very low density of gas molecules.

[<i>⊯Notes</i>	
		• Features:
		O Low Density: The exosphere contains very few air molecules, and collisions between molecules are rare.
		• Transition to Space: In the exosphere, atmospheric gases become extremely sparse, and individual gas particles can travel great distances before colliding with another particle. Eventually, these particles escape the Earth's gravitational pull and enter space.
		O No Well-Defined Boundary: The exosphere gradually fades into the emptiness of space, and there is no clear boundary between the Earth's atmosphere and space.
		• For Example: Satellites and space probes operate within the exosphere and beyond, orbiting the Earth or exploring other celestial bodies.
		In summary, the Earth's atmosphere consists of several layers, each with its distinct characteristics and functions. These layers interact and play a crucial role in shaping the planet's climate, weather patterns, and protecting life from harmful radiation. Understanding the structure and dynamics of the atmosphere is essential for various scientific disciplines, including meteorology, climatology, and space exploration.
		TEMPERATURE
		• Temperature is a relative measure or indication of hotness or coldness .
		 A hot utensil is said to have a high temperature, and an ice cube to have a low temperature.
		• An object with a higher temperature than another is said to be hotter.
		 Temperature is a measure of how fast the atoms in a material are vibrating.
		 High-temperature particles vibrate faster than low temperature particles. Danidh wibrating atoms cmash tegether, which generates heat
		 As a material cools down, the atoms vibrate more slowly and collide less frequently. As a result, they emit less heat.
		 What is the difference between heat and temperature? Heat is a form of energy while temperature denotes the intensity of hotness or coldness of any substance. Heat denotes the quantity of energy present in any substance while temperature refers to the degree of hotness of that substance.
		 Which has higher heat, and which has a higher temperature: a candle flame or a bathtub full of hot water? > The flame has a higher temperature, but less heat, because the hot region is very small.
		The bathtub has a lower temperature but contains much more heat because it has many more vibrating atoms. The bathtub has greater total energy.

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HEAT

- Heat is the form of energy transferred between two (or more) systems or a system and its surroundings by virtue of temperature difference.
- Heat is taken in or released when an object changes state, or changes from a gas to a liquid, or a liquid to a solid.
- This heat is called latent heat. When a substance changes state, latent heat is released or absorbed.
- A substance that is changing its state of matter does not change temperature.
- All of the energy that is released or absorbed goes toward changing the material's state.

Example:

Imagine a pot of boiling water on a stove burner: That water is at 100° C (212° F). If you increase the temperature of the burner, more heat enters the water. The water remains at its boiling temperature, but the additional energy goes into changing the water from liquid to gas. With more heat, the water evaporates more rapidly. When water changes from a liquid to a gas it takes in heat. Since evaporation takes in heat, this is called evaporative cooling. Evaporative cooling is an inexpensive way to cool homes in hot, dry areas.

The SI unit of heat energy transferred is expressed in joule (J) while the SI unit of temperature is Kelvin (K), and °C is a commonly used temperature unit.



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	LAPSE RATE
	• The atmospheric lapse rate refers to the change of an atmospheric variable with a change of altitude, the variable being temperature unless specified otherwise (such as pressure, density or humidity).
	 While usually applied to Earth's atmosphere, the concept of lapse rate can be extended to atmospheres (if any) that exist on other planets.
	 Lapse rates are usually expressed as the amount of temperature change associated with a specified amount of altitude change, such as 9.8 °Kelvin (K) per kilometer, 0.0098 °K per meter or the equivalent 5.4 °F per 1000 feet.
	If the atmospheric air cools with increasing altitude, the lapse rate may be expressed as a negative number. If the air heats with increasing altitude, the lapse rate may be expressed as a positive number.
	Types of lapse rates
	 There are <i>three types</i> of lapse rates that are used to express the rate of temperature change with a change in altitude, <i>namely</i>-
	1. The dry adiabatic lapse rate
	2. The wet adiabatic lapse rate
	3. The environmental lapse rate
	1. Dry adiabatic lapse rate
	 Atmospheric pressure decreases with altitude, causing the volume of an air parcel to expand as it rises.
	 Conversely, a sinking air parcel from a higher to a lower altitude experiences compression due to the higher pressure at the lower altitude.
	 Adiabatic lapse rate is the rate at which the temperature of an air parcel changes in response to expansion or compression during changes in altitude, assuming no heat exchange occurs during the process.
	 Saturated air: Earth's atmospheric air usually contains some water vapor. When it contains as much water vapor as it can hold, it's referred to as saturated air (relative humidity of 100%).
	 Dry adiabatic lapse rate: The lapse rate of unsaturated air (relative humidity less than 100%).
	 Also known as dry adiabatic, DALR, or unsaturated lapse rate.
	 The term "dry" indicates no liquid water (moisture) is present in the air, but water vapor may still be present.
	 Troposphere: The lowest layer of the Earth's atmosphere where almost all human activity takes place.
	 Dry adiabatic lapse rate is approximately constant in the troposphere due to little variation in gravity (g) and specific heat capacity of dry air (cpd) with altitude.
	2. Wet adiabatic lapse rate
	 An unsaturated parcel of air will rise from Earth's surface and cool at the dry adiabatic rate of – 9.8 K / kilometer until it reaches the atmospheric dew point.
	 At the dew point temperature, the air parcel becomes saturated, and the rate of cooling will decrease to the wet adiabatic lapse rate (SALR or MALR).

	<i>∝Notes</i>
 The wet adiabatic lapse rate is not constant and depends on the amount of water vapor the air contained when it started to rise, resulting in variable heat release. In the transporter the rate can yary from about 4 K (bilemeter (2.2.5) (1000 ft) at the transporter (2.2.5). 	
 In the troposphere, the rate can vary from about 4 K / kilometer (2.2 F / 1000 ft) at 25 °C (77 °F) to about 7 K / kilometer (3.8 °F / 1000 ft) at – 10 °C (14 °F). 	
 After reaching the dew point and cooling at the wet adiabatic lapse rate, the air parcel will continue to rise until all of its water vapor has condensed, and its cooling rate will revert back to the dry adiabatic lapse rate. 	
3. Environmental lapse rate	
 The dry adiabatic lapse rate and the wet adiabatic lapse rate are both theoretical rates. 	
• The actual real-world profile of temperature versus altitude that exists at any given time and in any given geographical location is called the environmental lapse rate, also often referred to as the ELR, prevailing lapse rate or ambient lapse rate.	
Altitude v/s Temperature	
• In Troposphere: The temperature in the troposphere generally decreases with altitude at an average rate of about 6.5°C per kilometer (3.6°F per 1,000 feet), known as the environmental lapse rate.	
• Cooling occurs because the troposphere is primarily heated from below. The Earth's surface is heated by the Sun , warming the air in contact with it.	
• Moving higher, the air encounters lower pressure , expanding and cooling, resulting in cooler temperatures at higher altitudes.	
 In Stratosphere: The temperature trend is different. The lower part of the stratosphere has a slight increase in temperature with altitude, while the upper part experiences a warming effect known as the stratopause. 	
• Warming is primarily due to the presence of the ozone layer , where ozone (O3) molecules absorb and scatter the Sun's ultraviolet (UV) radiation . This absorbed energy heats the surrounding air, causing the temperature to rise with increasing altitude.	
 In Mesosphere: Above the stratosphere is the mesosphere, where temperatures start to decrease again with altitude. The temperature reaches its coldest point at the mesopause, as low as -90°C (-130°F). 	
• Cooling occurs because there is less heat-absorbing gas compared to the stratosphere, and the density of air molecules decreases with altitude, resulting in a temperature decrease with height.	
• Thermosphere and Exosphere: In the thermosphere and exosphere, the relationship between altitude and temperature becomes more complex due to the presence of highly energetic particles from the Sun.	
• The thermosphere experiences a rapid increase in temperature due to the absorption of intense solar radiation.	
• The exosphere has very high temperatures for individual molecules despite an overall low density.	
• Overall, the altitude vs. temperature relationship in the Earth's atmosphere is crucial for understanding weather patterns, climate dynamics, and the behavior of various atmospheric layers. It influences many atmospheric processes, from the formation of clouds and precipitation in the troposphere to the absorption of harmful UV radiation by the ozone layer in the stratosphere.	

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	Inversion of temperature
	 Under normal conditions, temperature usually decreases with an increase in altitude in the troposphere at a rate of 1 degree for every 165 meters. This is called normal lapse rate.
	 Inversion of temperature in the atmosphere refers to a situation where the normal decrease in temperature with increasing altitude is reversed, leading to warmer air above cooler air. Inversions occur when a layer of warm air is trapped above a layer of cooler air, preventing the usual vertical mixing of the atmosphere. They can have significant effects on weather and air quality.
	Favorable Conditions for Temperature Inversion:
	 Favorable Conditions for Temperature Inversion: Long winter nights: Loss of heat by terrestrial radiation from the ground surface during the night may exceed the amount of incoming solar radiation. Cloudless and clear sky: Loss of heat through terrestrial radiation proceeds more rapidly without any obstruction. Dry air near the ground surface: It limits the absorption of the radiated heat from the Earth's surface. Slow movement of air: It results in no transfer or mixing of heat in the lower layers of the atmosphere. Snow-covered ground surface: It results in the maximum loss of heat through reflection of incoming solar radiation.
	Inversion) Temperature Inversion intermontane valley inversion
	Non-Advectional Temperature Inversion:
	 Advection refers to the horizontal movement of air masses, which can influence the temperature profile of a region. A non advectional temperature
	 A non-advectional temperature inversion occurs without the influence of horizontal air movement or advection. Other atmospheric processes are responsible for the formation and
	maintenance of the inversion layer.

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Types of Non-Advectional Temperature Inversion:

- 1. Radiation Inversion:
- **Radiation inversions** occur during clear, calm nights when the Earth's surface cools rapidly through **radiation**.
- The ground cools the air in direct contact with it, making the cooler air near the surface denser and causing it to stay close to the ground.
- Above this cool layer, there may be a layer of warmer air that acts as a cap, preventing further vertical mixing of the atmosphere and resulting in a **temperature inversion**.

Example: Formation of **fog** during a clear, calm night due to the cooling of the ground and air near the surface. The inversion layer traps the fog close to the ground, reducing visibility and creating potentially hazardous driving conditions.



- 2. Subsidence Inversion:
- Subsidence inversions occur when a large-scale sinking of air in the atmosphere leads to compressional heating.
- This typically happens in high-pressure systems when air descends from upper levels of the atmosphere to the surface.
- The sinking air gets compressed, causing its temperature to increase. This warmer air forms a stable layer above the cooler air near the surface, creating a subsidence inversion.



• **Example:** Common in some desert regions during high-pressure systems, where the sinking warm air creates a stable layer, trapping pollutants and dust near the surface. This can lead to the accumulation of smog and reduced air quality in these areas.

Impacts of Inversions:

- Inversions can have **positive** and **negative** impacts.
- **Positive**: They can protect crops from frost damage by trapping warm air near the surface on cold nights.
- **Negative:** When inversions trap pollutants close to the ground, they lead to poor air quality and increased health risks.

Adve	ctional Temperature Inversion:	Valley and mountain breezes
•	An advectional temperature inversion occurs due to the horizontal movement (advection) of warm air over a cooler surface or vice versa. Unlike radiation or subsidence inversions, which are primarily caused by vertical processes, advectional inversions result from horizontal temperature differences .	day warm air cool air warm air day warming mountainsides
1.	Valley Inversion in Intermontane Valley:	
•	Valley inversions typically occur during clear, calm nights when the Earth's surface cools rapidly through radiation .	night cooling mountainsides
•	The air near the surface cools, becoming colder than the air at higher elevations in the valley.	
•	Colder air near the surface is denser and set density compared to the warmer air above, o valley inversion.	tles into the valley due to its highe creating a stable layer known as th
Featu	ires and Impacts:	
1.	Temperature Gradient : The valley inversion gradient within the valley, with the air near the at higher elevations.	creates a noticeable temperature e valley floor significantly colder than
2.	Trapping of Air Pollutants : Stable conditions to near the surface, leading to poor air quality , significant sources of air pollution.	rap pollutants and particulate matte especially in urbanized valleys with
3.	Fog Formation: Colder air near the valley floc colder months, causing water vapor to conder	oor may reach its dew point during nse into fog or low-lying clouds.
4.	Impact on Local Weather: Valley inversions infl affecting temperature, humidity, and cloud of inhibit the heating of the valley floor during temperatures compared to surrounding moun	luence weather patterns in the valley cover. Trapped cold air can delay o the day, leading to cooler daytime tain slopes.
Va nights	lley inversions are common in regions with , making them prevalent in many intermontan	clear skies, light winds, and long e valleys worldwide.
2.	Frontal or Cyclonic Inversion:	
•	Fronts are boundaries where different air masses meet, and they can be warm fronts (warm air advancing over colder air) or cold fronts (cold air advancing over warmer air).	Frontal Inversion warmer, lighter air FRONTAL warm INVERSION cool cool rheavier air cool cool warm
•	As a front approaches, warm, less dense air is forced to rise over the cooler, denser air mass. As the warm air ascends, it cools and condenses, leading to cloud formation and often precipitation. This rising warm air associated with cyclones.	200 miles 100 miles Advancing Daker Cold Front
•	During the rising and cooling process, the surrounding air aloft, forming a temperature below the warmer air mass associated with th	air may become colder than th e inversion that traps the cooler a e front or cyclone.

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Features and Impacts:

- **1. Stability**: Frontal or cyclonic inversions create **stable atmospheric conditions**, preventing further vertical mixing of the air, resulting in calm and relatively wind-free conditions.
- 2. Weather Conditions: Frontal or cyclonic inversions influence the weather associated with the front or cyclone, enhancing cloud development and precipitation, leading to potentially heavy rain or snowfall.
- **3.** Impact on Air Quality: Like other temperature inversions, frontal or cyclonic inversions trap pollutants near the surface, leading to degraded air quality, especially in urban areas.
- **4. Dissipation:** Frontal or cyclonic inversions may dissipate as the front or cyclone moves through the region and the air masses mix, allowing the inversion to break down and the air to disperse.

WARM FRONTS VS COLD FRONTS

Aspect	Warm Front	Cold Front
Air Masses	Warm air mass advancing and rising over a retreating cold air mass	Cold air mass advancing and displacing a warm air mass
Weather Transition	Gradual and slow weather change	Rapid and abrupt weather change
Weather Conditions	Often associated with long- lasting, steady precipitation	Intense and short-lived precipitation, followed by clearing skies
Cloud Types	Often brings stratus and nimbostratus clouds	Often brings cumulonimbus clouds
Temperature Change	Temperature increases as the warm air replaces the cold air	Temperature drops as the cold air replaces the warm air
Pressure Change	Gradual decrease in atmospheric pressure	Rapid decrease in atmospheric pressure
Wind Shift	Winds shift from southerly to southwesterly ahead of the front	Winds shift from southerly to northwesterly behind the front
Visibility	Generally reduces due to prolonged precipitation	Briefly reduces due to heavy rain and thunderstorms
Examples	Rainy and overcast days with moderate temperatures	Thunderstorms and heavy rain, followed by cooler and drier conditions





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<i>∞Notes</i>	
	Distribution of Temperature in the Earth's Atmosphere:
	• The interaction of insolation (incoming solar radiation) with the atmosphere and the earth's surface creates heat , measured in terms of temperature .
	 Heat represents the molecular movement of particles comprising a substance, while temperature measures the intensity of heat.
	Heat and Temperature:
	 Temperature indicates the relative degree of heat of a substance.
	• Heat is the energy that makes things or objects hot, while temperature measures the average kinetic energy of the atoms in a substance.
	 Heat (energy) is the total kinetic energy of all the atoms in a substance; the more atoms present, the greater the heat.
	 Temperature represents the average kinetic energy of the atoms in a substance; a few atoms with rapid motion will have a higher temperature than many atoms with slow motion.
	 Although distinct from each other, heat and temperature are closely related because gain or loss of heat is necessary to raise or lower the temperature.
	• The movement of heat depends upon the temperature difference between two bodies.
	• Heat always moves from a body of higher temperature to that of lower temperature.
	 Temperature is primarily measured in the kelvin (K) unit in the study of physical sciences and most commonly measured in Celsius (C) or Fahrenheit (F) in day-to-day uses (denoted as °C, °F, and °K).
	Factors Affecting the Temperature Distribution: The temperature of the air at any place is influenced by the following factors:
	• The latitude of the place.
	• The altitude of the place.
	• Distance from the sea , the air-mass circulation .
	• The presence of warm and cold ocean currents .
	Local aspects.
	Global Temperature Distribution:
	 The global distribution of temperature can be understood by studying the temperature distribution in January and July.
	• The temperature distribution is generally shown on the map with the help of isotherms . Isotherms are lines joining places having equal temperature .
	• The effect of the latitude on temperature is well pronounced on the map, as the isotherms are generally parallel to the latitude .
	• The deviation from this general trend is more pronounced in January than in July , especially in the northern hemisphere.
	 In the northern hemisphere, the land surface area is much larger than in the southern hemisphere, leading to more pronounced effects of land mass and ocean currents.

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Seasonal Temperature Distribution – January



- During January, it is winter in the northern hemisphere and summer in the southern hemisphere.
- The **western margins** of continents are warmer than their eastern counterparts, since the **Westerlies** are able to carry high temperature into the landmasses.
- The temperature gradient is close to the **eastern margins** of continents. The isotherms exhibit a more regular behavior in the **southern hemisphere**.

Northern Hemisphere

- The isotherms deviate to the north over the ocean and to the south over the continent. This can be seen on the **North Atlantic Ocean**.
- The presence of warm ocean currents, **Gulf Stream** and **North Atlantic drift**, make the Northern Atlantic Ocean warmer, and the isotherms show a poleward shift indicating that the oceans are warmer and can carry high temperatures poleward.
- An equatorward bend of the isotherms over the northern continents shows that the landmasses are overcooled and that polar cold winds are able to penetrate southwards, even in the interiors. It is much pronounced in the **Siberian plain**.
- **Lowest temperatures** are recorded over northern Siberia and Greenland.

Southern Hemisphere

- The effect of the ocean is well pronounced in the southern hemisphere. Here the isotherms are more or less **parallel to the latitudes** and the variation in temperature is more gradual than in the northern hemisphere.
- The high-temperature belt runs in the southern hemisphere, somewhere along **30°S latitude**.
- The thermal equator lies to the south of the geographical equator (because the **Intertropical Convergence Zone or ITCZ** has shifted southwards with the apparent southward movement of the sun).

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Seasonal Temperature Distribution – July
 During July, it is summer in the northern hemisphere and winter in the souther hemisphere. The isothermal behavior is the opposite of what it is in January.
 In July, the isotherms generally run parallel to the latitudes. The equatorial ocear record warmer temperature, more than 27°C. Over the land, more than 30°C noticed in the subtropical continental region of Asia, along the 30° N latitude.
Northern Hemisphere
 The highest range of temperature is more than 60°C over the northeastern part of Eurasian continent. This is due to continentality. The least range of temperature 3°C, is found between 20°S and 15°N.
 Over the northern continents, a poleward bend of the isotherms indicates that the landmasses are overheated and the hot tropical winds are able to go far into the northern interiors.
 The isotherms over the northern oceans show an equatorward shift indicatir that the oceans are cooler and are able to carry the moderating effect into tropic interiors. The lowest temperatures are experienced over Greenland.
 The highest temperature belt runs through northern Africa, west Asia, northwere India, arid southeastern USA. The temperature gradient is irregular and follows zig-zag path over the northern hemisphere.
Southern Hemisphere
at c 30° c 10° c 20° c 20° c
Thermal Equator 0 Equator 0 C C C C C C C C C C C C C C C C C C C
20°C 10°C 10°C
60° 0 2000 6000 Km
90° 0° 90° 180°

CLIMATE CHANGE AND ATMOSPHERE

The interaction between the **lower and middle atmospheres** impacts the **upper atmosphere** primarily through the upward propagation of **atmospheric waves**. As these waves move upward, their magnitude increases due to the **exponential decrease** in **atmospheric density** with altitude.

Impacts of Climate Change on the Atmosphere:

- Heat Waves: Climate change will lead to the enlargement of stalled high-pressure weather systems known as "blocking events," which have already caused some of the deadliest heat waves of the 21st century.
- Contraction of the Upper Atmosphere: The rise in atmospheric CO2 concentration is causing cooling and compression of the upper atmosphere, resulting in lower densities at higher altitudes and consequently, a prolonged lifespan of space debris.
 - Additionally, there might be indirect consequences of climate change in the lower/middle atmosphere that influence long-term trends in upper atmosphere density.
- Formation of Surface-Level Ozone: Climate changes can affect local air quality by increasing ground-level ozone in many regions due to atmospheric warming.
- Uneven Warming and Cooling: Ozone in the atmosphere contributes to warming the climate, while different components of particulate matter (PM) can have warming or cooling effects on the climate.
 - For example, **black carbon**, a particulate pollutant from combustion, contributes to Earth's **warming**, whereas **particulate sulfates cool** the atmosphere.
- Drastic Climatic Changes: Climate models (IPCC, 2007) have been utilized to forecast the evolution of mean temperature and precipitation rate in the coming centuries.
 - Under a "business-as-usual" scenario, the projected global surface mean temperature increase by the end of the 21st century is 2.8°C, with an average warming of 3.5°C on land and potentially up to 7°C in the Arctic region.